

## **TE equation**

The equation used in the Excel spreadsheet ECT\*.xls\* from [www.thetropicalevents.com](http://www.thetropicalevents.com) to calculate the approximate time of a solstice or equinox, *TE*, is as follows:

$$eTE = aTE + \Delta bTE + \Delta mTE + \Delta eTE$$

*eTE* is an *eccentric* Tropical Event: a solstice or equinox calculated from pure ellipse, unperturbed by the Moon and planets.

*aTE* is an *average* Tropical Event: calculated from a circular orbit, with an *average* precession rate and an *average* sidereal mean motion. The initial calculations are in dynamical time, so that the quantity  $\Delta T$ , (dynamical time) – (Universal Time), needs to be included for studies within calendars. Use in a “Copernican” calendar includes an estimation of Universal Time.

$$aTE = Y \times aty$$

*aty* is the length of the *average* tropical year, calculated from the *average* sidereal year and the *average* precession rate. 1.0000010159 AU is used for the million year average Earth-Sun distance (Quinn, Tremaine, and Duncan 1991), so that an approximate value  $\sim 365.256362$  is used for the average sidereal year. 50.4712 arcseconds per Julian year is used for the average rate of precession (Laskar 1993), so that a value  $\sim 365.242138$  is used for *aty*. A good example is  $365^{77/318}$ , so that

$$aty = \frac{116147}{318}$$

*Y* is the number of average tropical years from the initial epoch  $Y_0$ . It rather simplifies the inclusion of an estimated  $\Delta T$  to set  $Y_0$  close to the epoch in which dynamical time = estimated Universal Time, and further simplifies calculations of solstices and equinoxes to set it as one of the average Tropical Events. We use the average Winter Solstice (northern hemisphere) *aWS* of 1819 as the initial epoch, as it is the closest Tropical Event to the *minimum* of the “best fit” parabola from Morrison and Stephenson 2004 for a long-term estimated  $\Delta T = -20 + 32t^2$  seconds, where  $t = (\text{year} - 1820)/100$ , so that

$$Y = \text{year} + \frac{Q}{4} - 1820$$

where *year* is the Gregorian calendar year. *Q* is the “quarter” year, where *Q* = 1 represents an average Vernal Equinox *aVE*; *Q* = 2 an average Summer Solstice *aSS*; *Q* = 3 an average Autumnal Equinox *aAE*; *Q* = 4 an average Winter Solstice *aWS*

As an example we will calculate an approximate time for the Vernal Equinox of 2010 in the Gregorian calendar

$$Y = 2010 + 1/4 - 1820 = 190.25$$

$$aTE = 190.25 \times 116147 / 318 = d\ 69487.3168239$$

*d* is the *day number* in dynamical time, counted from  $Y_0$  as 86400 seconds SI per *day*. It is the same time scale used in the familiar *Julian day numbers*, but shifted by the amount  $\Delta_{syst}$ , about 20 seconds SI, as the Julian day numbers are set so that dynamical time = Universal Time at approximately the year 1900, and we are using the approximate year 1820. It further simplifies a calendar construction to set the initial epoch  $Y_0$  at the beginning of a civil day in Greenwich, so that

$$JDE\ of\ Y_0 = 2385782.5 - \Delta_{syst} = JDE\ 2385782.499769$$

where  $\Delta_{syst} = 0.000231$  days, exactly 19.9584 seconds (20/86400 cannot be exactly represented as a decimal).

$$JDE\ of\ aTE = aTE + (JDE\ of\ Y_0) = 69487.3168239 + 2385782.499769 = JDE\ 2455269.8165929$$

This is March 14, 2010 at about 7:30 AM in Greenwich,  $6\frac{1}{2}$  days sooner than the true Vernal Equinox, but provides a good “first approximation” from which to calculate the three delta’s of the *TE* equation. Even for the 30 million years of Laskar’s La93(0,1) data this first approximation, the *aTE*, differs from the *eTE* by no more than  $\pm 20$  days of dynamical time *TD*. The *eTE* can differ from a *true* solstice or equinox by  $\pm 20$  minutes *TD*: more than adequate for long term calendar studies.

## $\Delta bTE$

$\Delta bTE$  is the least significant of the 3 delta's, but is the most elusive, as it requires integral knowledge of the sidereal mean motion. The only data readily available is from Laskar (1986) given in terms of  $t$ , units of 10000 Julian years from J2000, for the (heliocentric) mean longitude of Earth  $\times 10^{10}$ , in radians, from the fixed Vernal Equinox of J2000:

$$17534703144 + 628307584918000t - 9793168t^2 + 429738t^3 + 734935t^4 + 83525t^5 - 59447t^6 - 52555t^7 + 13798t^8 + 14426t^9 - 564t^{10} \quad (1)$$

and is valid for  $\pm 10000$  Julian years from J2000

The derivative of this formula reveals the sidereal mean motion

For epochs outside the range  $-10000$  to  $+10000$  the sidereal year can simply be the same as the average sidereal year  $asy$ , and if one is using  $365^{77}/_{318}$  for the average tropical year  $aty$ , for consistency with the average precession of La93(0,1),  $50''.4712$  per Julian year, one should use a value close to  $365.256362539$ , and  $365^{413}/_{1611}$  works just fine, so that

$$asy = \frac{588428}{1611}$$

The average speed of the precession during the average sidereal year  $asy$  is indicated by  $asp_{asy}$

$$a_1 = \text{den}_{asy} \times \text{num}_{aty} = 1611 \times 116147 = 187112817$$

$$b_1 = a_1 - \text{den}_{aty} \times \text{num}_{asy} = 187112817 - 318 \times 588428 = 7287$$

$\text{gcf}_1 = \text{gcf}(a_1, b_1) = 3$ , the greatest common factor of  $a_1$  and  $b_1$

$$\text{the constant } a = a_1 / \text{gcf}_1 = 187112817 / 3 = 62370939$$

$$\text{the constant } b = b_1 / \text{gcf}_1 = 7287 / 3 = 2429$$

$$asp_{asy} = a / b = 25677.62000823384108\dots$$

$asp_{Jy}$  indicates the average speed of the precession during a Julian year,  $365.25$  days

$asp_{Jy} = asp_{asy} \times asy / 365.25 = \frac{a \times 588428 \times 4}{b \times 1611 \times 1461} = \frac{146803227575568}{5717066859} = \frac{273376587664}{10646307} = 25678.0673020231\dots$ , slightly large, showing  $50''.47108$  per Julian year, but quite good enough for approximate times of solstices and equinoxes for the 30 million years of Laskar's data. If one used  $asy = 365^{5097}/_{19882}$  it would show  $50''.471199$ , closer to the recommended value of  $50''.4712$

The constant  $\Delta bTE_{J2000} = -0.13$  days

$$\sim aTE = JDE \text{ of } aTE - 2451545 = 2455269.8165929 - 2451545 = 3724.8165929$$

$$(\text{avg trop yrs from J2000}) = \sim aTE / aty = 3724.8165929 \times 318 / 116147 = 10.1982115469379$$

$T_{aTE}$  is the time T in Julian millennia from J2000, at the time of the  $aTE$

$$T_{aTE} = \sim aTE / 365250 = 3724.8165929 / 365250 = 0.010197992040794$$

Using Laskar 1986 formula (1)

$$(\text{radians from J2000}) = \frac{628307584918000t - 9793168t^2 + 429738t^3 + 734935t^4 + 83525t^5 - 59447t^6 - 52555t^7 + 13798t^8 + 14426t^9 - 564t^{10}}{10^{10}}$$

$$\text{where } t = T_{aTE}/10 = 0.0010197992040794$$

$$(\text{radians from J2000}) = 64.0747575006$$

$$(\text{sid yrs from J2000}) = (\text{radians from J2000}) / 2\pi = 64.0747575006 / 2\pi = 10.197814383635$$

asy(T) is the average length of the sidereal year during time span T, where T is Julian millennia from J2000

$$\text{asy}(T) = \sim aTE / (\text{sid yrs from J2000}) = 3724.8165929 / 10.197814383635 = 365.256363057307$$

$$A_1 = a \times \text{num}_{\text{asy}} = 62370939 \times 588428 = 36700806893892$$

$$B_1 = b \times \text{den}_{\text{asy}} = 2429 \times 1611 = 3913119$$

gcf<sub>2</sub> = gcf(A<sub>1</sub>, B<sub>1</sub>) = 537, the greatest common factor of A<sub>1</sub> and B<sub>1</sub>

$$\text{the constant A} = A_1 / \text{gcf}_2 = 36700806893892 / 537 = 68344146916$$

$$\text{the constant B} = B_1 / \text{gcf}_2 = 3913119 / 537 = 7287$$

$$\text{avg bty}(T) = \text{asy}(T) / (1 + B \times \text{asy}(T) / A)$$

$$\text{avg bty}(T) = 365.256363057307 / (1 + 7287 \times 365.256363057307 / 68344146916) = 365.242138914285$$

$$\sim bTE = (\text{avg bty}(T)) \times (\text{avg trop yrs from J2000}) = 365.242138914285 \times 10.1982115469379 = 3724.81659850399$$

$$\sim \Delta bTE = \sim bTE - \sim aTE = 3724.81659850399 - 3724.8165929 = 0.00000560399$$

$$\Delta bTE = \sim \Delta bTE + \Delta bTE_{J2000} = 0.00000560399 - 0.13 = -0.12999439601 \text{ days}$$

Outside the range  $-10000$  to  $+10000$  one can simply set  $\Delta bTE = 0$ , and then add  $\Delta bTE_{J2000}$  to  $\Delta mTE_{y0}$  in the following section. The additional error in calculating a solstice or equinox *without* including  $\Delta bTE$  from  $-10000$  to  $+10000$  can be  $\pm 1\frac{1}{2}$  hours  $TD$ . For an interesting study of *possible* long term effects due to the variations of amplitudes (*possibly*  $\pm 5$  hours) use (with caution) the “fictitious rendition” for the length of the sidereal year from ECT\*.xls\*

### $\Delta mTE$

$\Delta mTE$  is a value calculated from the precession pA. It can be calculated by the following method for the entire range of La93(0,1) data, with excellent results by the use of Lagrange interpolation polynomials (see Meeus 1998, or documentation for the “Orbital Elements” Excel add-in at [thetropicalevents.com](http://thetropicalevents.com)). For this example we use formulae from Laskar 1986

(pA from J2000), valid for  $\pm 10000$  Julian years from J2000:

$$(502909.66t + 11119.71t^2 + 77.32t^3 - 2353.16t^4 - 180.55t^5 + 174.51t^6 + 130.95t^7 + 24.24t^8 - 47.59t^9 - 8.66t^{10}) \times \pi / (180 \times 60 \times 60) \quad (2)$$

$$\text{where } t = T_{aTE} / 10 = 0.0010197992040794$$

$$(\text{pA from J2000}) = 0.00248650482258583 \text{ radians}$$

The following if statement converts various data between  $-3\pi$  and  $+3\pi$  to a value between  $-\pi$  and  $+\pi$

$$(\text{pA} \pm \pi) = \text{IF}((\text{pA from J2000}) > \pi, (\text{pA from J2000}) - 2\pi, \text{IF}((\text{pA from J2000}) <= -\pi, (\text{pA from J2000}) + 2\pi, (\text{pA from J2000}))) \\ = 0.00248650482258583$$

$$(\text{pA}/2\pi \text{ circle}) = \text{IF}((\text{pA} \pm \pi) < 0, (\text{pA} \pm \pi) / 2\pi + 1, (\text{pA} \pm \pi) / 2\pi) = 0.000395739533536371$$

$$(\text{total avg cycles}) = \sim aTE / (365.25 \times \text{asp}_{Jy}) = \sim aTE \times b / (\text{asy} \times a) = 3724.8165929 \times 2429 \times 1611 / (62370939 \times 588428) \\ = 0.000397147959807337$$

$$(\sim \text{cycles} \pm 1 \text{ from J2000}) = \text{INT}(\text{total avg cycles}) + (\text{pA}/2\pi \text{ circle}) = 0.000395739533536371$$

where INT(x) is the integer less than or equal to x

$$(T_{aTE} \text{ at } Y_0) = ((JDE \text{ of } Y_0) - 2451545) / 365250 = -0.180047913021219$$

(pA at  $Y_0$ ): using Laskar 1986 precession, formula (2), with  $t = (T_{aTE} \text{ at } Y_0)/10 = -0.0180047913021219$

$$(pA \text{ at } Y_0) = -0.0438813563532027 \text{ radians}$$

$$\begin{aligned} (\text{cycles } Y_0 \text{ to aTE}) &= (\sim\text{cycles} \pm 1 \text{ from J2000}) - \text{ROUND}((\sim\text{cycles} \pm 1 \text{ from J2000}) - (\text{total avg cycles}), 0) - (pA \text{ at } Y_0) / 2\pi \\ &= 0.000395739533536371 - 0 + 0.0438813563532027 / 2\pi = 0.007379674307 \end{aligned}$$

where ROUND( $x,0$ ) rounds to the integer closest in value to  $x$

$$(M\text{sidereal orbits}) = Y \times aty / asy = Y \times a / (a + b) = 190.25 \times 62370939 / (62370939 + 2429) = 190.242591112765$$

$$\begin{aligned} (\text{total mean tropical years from } Y_0 \text{ to aTE}) &= (M\text{sidereal orbits}) + (\text{cycles } Y_0 \text{ to aTE}) \\ &= 190.242591112765 + 0.007379674307 = 190.249970787072 \end{aligned}$$

$$\begin{aligned} \sim mTE &= Y \times aTE / (\text{total mean tropical years from } Y_0 \text{ to aTE}) = 190.25 \times 69487.3168239 / 190.249970787072 \\ &= 69487.3274936939 \end{aligned}$$

The constant  $\Delta mTE_{Y_0} = 8.422 \text{ days}$ , given to the nearest  $1/1000$  of a day, indicating that the intended accuracy at this point in the process is  $\pm 0.0005 \text{ days}$ , or about 1 *minute TD*.

$$\Delta mTE = \sim mTE - aTE + \Delta mTE_{Y_0} = 69487.3274936939 - 69487.3168239 + 8.422 = 8.4326697939 \text{ days}$$

$\Delta mTE$  varies by about  $\pm 12 \text{ days}$  during the entire 30 million years of Laskar's La93(0,1) data

$$mTE = aTE + \Delta bTE + \Delta mTE$$

$$mTE = aTE + \Delta bTE + \Delta mTE = 69487.3168239 - 0.12999439601 + 8.4326697939 = d 69495.6194992979$$

$$JDE \text{ of } mTE = mTE + (JDE \text{ of } Y_0)$$

$$JDE \text{ of } mTE = 69495.6194992979 + 2385782.499769 = JDE 2455278.1192683$$

This is March 22, 2010 at about 3 PM in Greenwich, almost 2 days later than the true Vernal Equinox.

The *mean* Tropical Event  $mTE$  represents the mean planet (as if we were in circular orbit) from the moving Vernal Equinox (including the precession), so that the longitude of the mean planet at the time  $T_{mTE}$ ,

$$T_{mTE} = (JDE \text{ of } mTE - 2451545) / 365250 = (2455278.1192683 - 2451545) / 365250 = 0.0102207235271732,$$

is a  $90^\circ$  increment (a solstice or equinox), with an intended accuracy of  $\pm 0.00001$  radians (about 1 minute of time). It should be noted that the constant  $\Delta mTE_{Y_0}$  is intended to include the "constant of aberration" for the "observed" Sun, about 0.000099 radians. Using  $T_{mTE}/10$  in equations (1) and (2) results in  $65.9710514681048 + 0.00249204740661573 - 0.000099 = 65.9734445155114$ , which reduces to 3.14159144371554 radians, 0.00000120987425 less than  $\pi$  (heliocentric longitude of the Vernal Equinox), well within our intended accuracy. Note: alternately one could skip the previous two sections by instead using an iterative scheme to find the time  $T_{mTE}$  of the mean planet at the required  $90^\circ$  increment, and then proceed as follows.

### $\Delta eTE$

$\Delta eTE$  is a generic term for each of the 4 values associated with each of the 4 Tropical Events, and is calculated from the eccentricity  $e$  and perihelion from the moving equinox  $\varpi$ , including the indication of the speed of the precession  $sp$  and the sidereal year  $sy$ , so that we can use the mean tropical year  $mty$ . The following method also applies to the use of Laskar's data for  $k$  and  $h$  to determine  $e$  and  $\omega$ , and the derivatives of precession, again best by using Lagrange interpolation polynomials.

$dpA$ : using the derivative of the precession in radians (Laskar 1986) per 1000 Julian years

$$dpA = (502909.66 + 22239.42t + 231.96t^2 - 9412.64t^3 - 902.75t^4 + 1047.06t^5 + 916.65t^6 + 193.92t^7 - 428.31t^8 - 86.6t^9) \times \pi / (10 \times 180 \times 60 \times 60)$$

$$\text{with } t = T_{mTE}/10 = 0.00102207235271732$$

$$dpA = 0.243828503601382 \text{ radians per 1000 Julian years}$$

Using the derivative of mean longitude of Earth from Laskar 1986 to determine the length of the sidereal year  $sy$

$$sy = \frac{2\pi \times 3652500 \times 10^{10}}{628307584918000 - 19586336t + 1289214t^2 + 2939740t^3 + 417625t^4 - 356682t^5 - 367885t^6 + 110384t^7 + 129384t^8 - 5640t^9}$$

$$\text{with } t = T_{mTE}/10 = 0.00102207235271732$$

$$sy = 365.256363062974$$

$$sp_{sy} = 2\pi \times 365250/(dpA \times sy) = 2\pi \times 365250/(0.243828503601382 \times 365.256363062974) = 25768.4222975507$$

$$mty = sy \times sp_{sy} / (sp_{sy} + 1) = 365.256363062974 \times 25768.4222975507 / 25769.4222975507 = 365.2421890408$$

Next we need the eccentricity  $e$  and perihelion in radians from the fixed equinox of J2000  $\omega$ , calculated from  $k$  and  $h$  Using Laskar 1986, again valid for  $\pm 10000$  Julian years from J2000:

$$k = \frac{-37408165 - 82266699t + 27626329t^2 + 11695572t^3 - 2695722t^4 - 715070t^5 + 218146t^6 + 22635t^7 - 19921t^8 - 2032t^9 + 475t^{10}}{10^{10}}$$

$$h = \frac{162844766 - 62030259t - 33829810t^2 + 8510121t^3 + 2770542t^4 - 467407t^5 - 62395t^6 + 247t^7 + 403t^8 + 686t^9 - 423t^{10}}{10^{10}}$$

$$\text{with } t = T_{mTE}/10 = 0.00102207235271732$$

$$k = -0.00374922186467685$$

$$h = 0.0162781331256635$$

$$e = \text{SQRT}(k^2 + h^2) = 0.0167043192811738, \text{ where } \text{SQRT}(x) \text{ is } \sqrt{x}$$

$$\omega = \text{ATAN2}(k, h) = 1.79717108039395, \text{ where } \text{ATAN2}(k, h) \text{ is the inverse tangent of } h/k, \text{ between } -\pi \text{ and } +\pi \text{ radians}$$

$$\varpi_1 = \omega + (pA \pm \pi) = 1.79717108039395 + 0.00248650482258583 = 1.79965758521654$$

note: here it is more accurate to recalculate  $(pA \pm \pi)$  using  $T_{mTE}$ , but the difference is minimal compared to the accuracy of this section:  $\pm 20$  minutes  $TD$ , due to the perturbations of the Moon ( $\pm 10$  mins) and of the planets ( $\pm 10$  mins).

converted to an angle between 0 and  $2\pi$  for perihelion from the moving equinox  $\varpi$

$$\varpi = \text{IF}(\varpi_1 > 2\pi, \varpi_1 - 2\pi, \text{IF}(\varpi_1 < 0, \varpi_1 + 2\pi, \varpi_1)) = 1.79965758521654$$

Now we use the “equation of center”  $C$ . From Meeus 1998 p.275 and the “Equation of Kepler” we find the true anomaly  $v$ , the eccentric anomaly  $E$ , and the mean anomaly  $M$

$$v \text{ at } VE = \pi - \varpi$$

$$v \text{ at } SS = 3\pi/2 - \varpi$$

$$v \text{ at } AE = 2\pi - \varpi$$

$$v \text{ at } WS = \pi/2 - \varpi$$

$$\tan \frac{E}{2} = \sqrt{\frac{1-e}{1+e}} \tan \frac{v}{2}$$

$$M = E - e \sin E$$

$$-C = M - v$$

$$n_t = \frac{2\pi}{mty}$$

$$\Delta eTE = \frac{-C}{n_t}$$

$$v \text{ at } VE = \pi - \varpi = 3.14159265358979 - 1.79965758521654 = 1.34193506837325$$

$$E = \text{ATAN}(\text{SQRT}((1-e)/(1+e)) \times \text{TAN}(v/2)) \times 2 = 1.32569630256058$$

$$M = E - e \sin E = 1.32569630256058 - 0.0167043192811738 \times \text{SIN}(1.32569630256058) = 1.30949122429074$$

$$-C = M - v = 1.30949122429074 - 1.34193506837325 = -0.0324438440825101$$

$$n_t = 2\pi / \text{mty} = 6.28318530717959 / 365.2421890408 = 0.0172027917248019$$

$$\Delta eTE = -C/n_t = -0.0324438440825101 / 0.0172027917248019 = -1.885963895 \text{ days}$$

Alternately, by calculating the lengths of the seasons, one can use the ‘‘Ancient Astronomer’s Formula’’ to calculate  $\Delta eTE$

$$\Delta eVE = \frac{F + 3W - 3V - S}{8}$$

$$\Delta eSS = \frac{W + 3V - 3S - F}{8}$$

$$\Delta eAE = \frac{V + 3S - 3F - W}{8}$$

$$\Delta eWS = \frac{S + 3F - 3W - V}{8}$$

$$E_{WS} = \text{ATAN}(\text{SQRT}((1-e)/(1+e)) \times \text{TAN}((\pi/2 - \varpi)/2)) \times 2 = -0.225101894822176$$

$$\varpi \text{ to } WS = (E_{WS} - e \times \text{SIN}(E_{WS})) \times \text{mty} / 2\pi = -12.8684575907845$$

$$E_{VE} = \text{ATAN}(\text{SQRT}((1-e)/(1+e)) \times \text{TAN}((\pi - \varpi)/2)) \times 2 = 1.32569630256058$$

$$\varpi \text{ to } VE = (E_{VE} - e \times \text{SIN}(E_{VE})) \times \text{mty} / 2\pi = 76.1208555703667$$

$$E_{SS} = \text{ATAN}(\text{SQRT}((1-e)/(1+e)) \times \text{TAN}((3\pi/2 - \varpi)/2)) \times 2 = 2.90891036561448$$

$$\varpi \text{ to } SS = (E_{SS} - e \times \text{SIN}(E_{SS})) \times \text{mty} / 2\pi = 168.87134312852$$

$$E_{AE} = \text{ATAN}(\text{SQRT}((1-e)/(1+e)) \times \text{TAN}((2\pi - \varpi)/2)) \times 2 = -1.78335716116808$$

$$\varpi \text{ to } AE = (E_{AE} - e \times \text{SIN}(E_{AE})) \times \text{mty} / 2\pi = -102.717560011001$$

$$W_1 = (\varpi \text{ to } VE) - (\varpi \text{ to } WS) = 88.9893131611513$$

$$V_1 = (\varpi \text{ to } SS) - (\varpi \text{ to } VE) = 92.7504875581535$$

$$S_1 = (\varpi \text{ to } AE) - (\varpi \text{ to } SS) = -271.588903139521$$

$$F_1 = (\varpi \text{ to } WS) - (\varpi \text{ to } AE) = 89.8491024202163$$

$$W = \text{IF}(W_1 < 0, W_1 + \text{mty}, W_1) = 88.9893131610537 \text{ days of Winter}$$

$$V = \text{IF}(V_1 < 0, V_1 + \text{mty}, V_1) = 92.7504875581535 \text{ days of Spring}$$

$$S = \text{IF}(S_1 < 0, S_1 + \text{mty}, S_1) = -271.588903139521 + 365.2421890408 = 93.653285901279 \text{ days of Summer}$$

$$F = \text{IF}(F_1 < 0, F_1 + \text{mty}, F_1) = 89.8491024202163 \text{ days of Fall}$$

$$\Delta eVE = (F + 3 \times W - 3 \times V - S) / 8 = -1.885963334 \text{ days}$$

The two methods calculate differently by 0.00000056 *days* (0.048 *seconds*), and one assumes that this is due to “decimal rounding”, and the “double precision” used by Excel, and most computer software.

$$eTE = mTE + \Delta eTE$$

$$eVE = mTE + \Delta eVE = 69495.6194992979 - 1.885963334 = d\ 69493.7335359639$$

$$JDE\ of\ eVE = eVE + (JDE\ of\ Y_0) = 69493.7335359639 + 2385782.499769 = JDE\ 2455276.23330496$$

## $\Delta T$

$\Delta T$ , the difference (dynamical time) – (Universal Time), is by far the most elusive quantity for long term calendar studies. However, using the study of Morrison and Stephenson (2004), Espenak and Meeus have developed a great set of equations for the years –1999 to +3000 available at <http://eclipse.gsfc.nasa.gov/SEcat5/deltatpoly.html> . We prefer to use them only for the years –404.15 to +2003.45, as the following “sum of 4 sines” more accurately follows the Morrison and Stephenson (2004) study from –500 to –2000, and allows an extension even further. This equation for  $\Delta T$  in seconds is

$$\begin{aligned} & -3.0169675 + 0.003390245877 * Y^2 \\ & + (11.85034251 * \sin(4.521017826 + 0.00009728265802 * Y) / 0.00009728265802 + 119589.730883314) \\ & + (4.889524586 * \sin(1.038218036 + 0.0003290065396 * Y) / 0.0003290065396 - 12803.1780446892) \\ & + (1.416055354 * \sin(3.116643354 + 0.004835133099 * Y) / 0.004835133099 - 7.30609135253061) \\ & + (1.209213516 * \sin(0.4888837632 + 0.004072065294 * Y) / 0.004072065294 - 139.46138984887) \end{aligned}$$

where Y can be  $d/aty$  , and is valid from  $Y = -7727.87259149758$  to  $Y = 1165.47962600512$  (5908 BC to 2985 AD)

To facilitate use of this formula unassociated with the above “method”, calculate Y as  $(JDE - 2385782.5) / 365.2421378$   
It should be noted by the use of "JDE" that this is intended to convert dynamical time into Universal time.

This single equation is no more than 15 seconds in error for the entire time span for the recorded values of  $\Delta T$ , ~1600 to 2010. One must include the errors associated with these recorded values. For example, the year 1600 is often given as 120 seconds, with an associated error of  $\pm 20$  seconds. A value of 85 seconds for the year 1600 would therefore be only 15 seconds in error. This equation returns 100.198 secs for the year 1600, so for that particular year it can be considered as having no error.

For the Vernal Equinox of 2010

$$Y = d / aty = 69493.7335359639 \times 318 / 116147 = 190.267568378318 \quad , \quad \text{and the “sum of 4 sines” returns}$$

$$\Delta T = 66.1405146650131 \text{ seconds}$$

However, it is often more satisfying to use the *extreme accuracy* of the Espenak and Meeus polynomials, where they use  $t = y - 2000$  , and  $y$  is given as:  $year + (month - 0.5) / 12$ . By this method  $y$  calculates the portion of the year at “mid-month” only. To facilitate smooth transitions we prefer to use:  $y = 1820 + (d - 18) / 365.2425$ . When not using the above “method” we recommend calculating  $y$  as  $(JDE - 2451544.5) / 365.2425 + 2000$  to obtain smooth transitions for their polynomials.

In order to transition from the Espenak and Meeus polynomials at 2003.45 to the “sum of sines” at 2050 we use a polynomial that also follows the data from the *International Earth Rotation and Reference Systems Service* ([IERS](http://www.iers.org)) from 2000 thru 2010:

$$63.9 + 0.164954 \times t - 0.00281933 \times t^2 + 0.000879724 \times t^3 - 0.0000104809 \times t^4 \quad \text{where } t = (JDE - 2451544.5) / 365.2425$$

$$t = (2455276.23330496 - 2451544.5) / 365.2425 = 10.2171387638629 \quad , \quad \text{so that the polynomial returns}$$

$$\Delta T = 66.1151192185045 \text{ seconds} \quad , \quad \text{and dividing by } 86400: \Delta T_{JD_{\text{sys}}} = 0.00076522 \text{ days}$$

$$JD\ of\ eVE = JDE\ of\ eVE - \Delta T_{JD_{\text{sys}}} = 2455276.23330496 - 0.00076522 = JD\ 2455276.23253974$$

$$\Delta T_{D_{\text{sys}}} = \Delta T_{JD_{\text{sys}}} + \Delta_{\text{sys}} = 0.00076522 + 0.000231 = 0.00099622 \text{ days}$$

$$D\ of\ eVE = eVE - \Delta T_{D_{\text{sys}}} = 69493.73353596 - 0.00099622 = D\ 69493.73253974$$

This is March 20, 2010 at 5:35 PM in Greenwich, 4 minutes after the true Vernal Equinox at 5:31 PM (calculated from Bretagnon and Simon 1986, including their formulae for nutation and aberration). This does *not* include the concept of “daylight savings time”.

The following formulae return ΔT in days from “Scenario 10a” in ECT1.6.xlsm, readjusted for standard double precision on an Excel spreadsheet, and Y is (JDE-2385782.5)/365.2421378.

From Y = -18080.8569219084 to Y = -7727.87259149758 (16261 BC to 5908 BC):

$$(-152.934868+0.003390245877*Y^2 + (11.85034251*\text{SIN}(4.521017826+0.00009728265802*Y)/0.00009728265802+119589.730883314) + (4.889524586*\text{SIN}(1.038218036+0.0003290065396*Y)/0.0003290065396-12803.1780446892))/86400 + 0.000002930134880633*Y + 0.0231025306199516$$

From Y = -1938856065.12609 to Y = -18080.8569219084:

$$-(1808.41540551846*Y-17802587981659.3*((1+(Y+18080.8569219084)/13463985755.3156)^(35/26)-1) - 69715684935.6374*((1+(Y+18080.8569219084)/13463985755.3156)^(11/2)-1)+32697685.4793694)+1.30023618539E-8$$

From Y = -4500000000 to Y = -1938856065.12609:

$$-(1808.41540551846*Y-4011240664619.57*((1+(Y+1938856065.12609)/2561143949.12143)^(14/13)-1) - 18115544946.9063*((1+(Y+1938856065.12609)/2561143949.12143)^2-1)+3402226352645.89)-0.000335693$$

From Y = +1165.47962600512 to Y = +10682.4663136617 (2985 AD to 12502 AD):

$$(-152.934868+0.003390245877*Y^2 + (11.85034251*\text{SIN}(4.521017826+0.00009728265802*Y)/0.00009728265802+119589.730883314) + (4.889524586*\text{SIN}(1.038218036+0.0003290065396*Y)/0.0003290065396-12803.1780446892))/86400 + 0.00000583377118476*Y + 0.00762862609185111 - 6.9389E-18$$

From Y = +10682.4663136617 to Y = +2500000000:

$$-(1808.4159809978*Y-1810574188741.9*((1+Y/1070750140.18301)^(20/19)-1) - 18105741887.419*((1+Y/1070750140.18301)^(32/19)-1)-0.798712707645782)$$

From Y = +2500000000 to Y = +5000000000:

$$-(1808.4159809978*Y-1672674863864.45*((1+(Y-2500000000)/965163563.979899)^(27/26)-1) - 21153783485.5091*((1+(Y-2500000000)/965163563.979899)^(3/2)-1)-455212309211.431)-0.00058841706$$

For an interesting study of different possible long term effects of ΔT use some of the other scenarios in ECT\*.xls\*

A single equation for ΔT in seconds from Y = -331.298643742993 to Y = 189.864878763783 (AD 1488 to 2009.813):

$$12.34471667+0.003390245877*Y^2 + (11.85034251*\text{Sin}(4.521017826+0.00009728265802*Y)/0.00009728265802+119589.730883314) + (4.889524586*\text{Sin}(1.038218036+0.0003290065396*Y)/0.0003290065396-12803.1780446892) + (1.416055354*\text{Sin}(3.116643354+0.004835133099*Y)/0.004835133099-7.30609135253061) + (1.209213516*\text{Sin}(0.4888837632+0.004072065294*Y)/0.004072065294-139.46138984887) + (0.348873982*\text{Sin}(2.711947318+0.03080982016*Y)/0.03080982016-4.71677204529075) + (0.2198280214*\text{Sin}(0.6000046074+0.01828535077*Y)/0.01828535077-6.7882249137094) + (0.4205366394*\text{Sin}(5.173509492+0.09238356146*Y)/0.09238356146+4.07662865593031) + (0.181331671*\text{Sin}(6.848632475+0.06187173456*Y)/0.06187173456-1.57028600582809) + (0.3309858275*\text{Sin}(2.696813396+0.1010673329*Y)/0.1010673329-1.40905575588025) + (0.2713170885*\text{Sin}(1.758701847+0.1327208509*Y)/0.1327208509-2.00828480358683) + (0.175825584*\text{Sin}(0.7598732321+0.1997702484*Y)/0.1997702484-0.60626574534795) + (0.1163835709*\text{Sin}(3.160230197+0.2762080472*Y)/0.2762080472+7.85269764264254E-3) + (0.09261707963*\text{Sin}(0.6334367883+0.2793572382*Y)/0.2793572382-0.196242431358713)$$

Better yet use until Y = 188.614438520988 (AD 2008.563), and then use the polynomial until Y = 230.05 (AD 2050)

For impeccable mathematical accuracy in Excel use the Xnumbers add-in, [from the author](#): Leonardo Volpi, or from an extended version supporting updates since the release of Excel 2007 available at [thetropicalevents.com](http://thetropicalevents.com)



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For questions or comments email to: [steve@thetropicalevents.com](mailto:steve@thetropicalevents.com)